Energy transfers in turbulence under rotation Some experiments



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and

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Hydrodynamics under rotation

Geo and astrophysical flows



Earth rotation from Galileo (1990)



Hurricane Irene (2011)



Pierre-Philippe Cortet

Hydrodynamics under rotation

Geo and astrophysical flows

- Geophysics (ocean, atmosphere, liquid core)
- Astrophysics

(gazeous planets, galaxies)







Physical ingredients

- Stratification → Internal gravity waves
- Magnetism → Alvén waves, Dynamo action
- Thermal convection
- <u>Rotation → Internal inertial waves</u>
- <u>Turbulence</u>

Model experiments



Hydrodynamics under rotation



In a non-dimensional version

$$Ro_t \frac{\partial \vec{u}}{\partial t} + Ro(\vec{u} \cdot \vec{\nabla})\vec{u} = -\vec{\nabla}p - \vec{e_z} \times \vec{u} + Ro Re^{-1}\Delta \vec{u}$$

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Turbulence under rotation

Navier-Stokes equation in a rotating frame

$$\vec{\Omega} \quad \underbrace{\frac{\partial \vec{u}}{\partial t}}_{\frac{\partial t}{\partial t}} + (\vec{u} \cdot \vec{\nabla})\vec{u} = -\frac{1}{\rho}\vec{\nabla}p - 2\vec{\Omega} \times \vec{u} + \nu\Delta\vec{u}$$

Reynolds number $Re = \frac{|(\vec{u} \cdot \vec{\nabla})\vec{u}|}{|\nu \Delta \vec{u}|} \sim \frac{UL_{\perp}}{\nu} \gg 1 \Rightarrow$ The flow is turbulent. Rossby number $Ro = \frac{|(\vec{u} \cdot \vec{\nabla})\vec{u}|}{|2\vec{\Omega} \times \vec{u}|} \sim \frac{U}{2\Omega L_{\perp}} \lesssim 1 \Rightarrow$ Rotation interacts with non-linearities. Temporal Rossby number $Ro_t = \frac{|\partial \vec{u}/\partial t|}{|2\vec{\Omega} \times \vec{u}|} \sim \frac{\sigma}{2\Omega} \simeq Ro$

This interaction drives turbulence toward an anisotropic 2D states invariant along $\vec{\Omega}$

Two-dimensionalisation of turbulence by rotation



Hopfinger, Browand & Gagne, JFM (1982) Oscillated grid in a rotating tank Emergence of columnar vortices aligned with the rotation axis



Jacquin, Leuchter, Cambon & Mathieu, JFM (1990)

Hot wire measurements in a rotating grid turbulence in a wind tunnel

Growth of axial integral scale with Ω



Rotation drives turbulence toward a 2D state

$$Ro_t \frac{\partial \vec{u}}{\partial t} + Ro(\vec{u} \cdot \vec{\nabla})\vec{u} = -\vec{\nabla}p - \vec{e_z} \times \vec{u} + Ro Re^{-1}\Delta \vec{u}$$

1. In the limit $Ro = Ro_t = 0$, NS becomes $\frac{1}{\rho} \vec{\nabla} p = -2\vec{\Omega} \times \vec{u}$ Taking its curl gives $(\vec{\Omega} \cdot \vec{\nabla})\vec{u} = \vec{0}$ Taylor-Proudman Theorem = Geostrophic equilibrium $Ro = 0 \rightarrow 2D$ 3C flow, but no turbulence



2D-3C flow

2. Whatever *Ro* and *Re*, for a pure 2D 3C flow, i.e. $\vec{u} = (\partial_y \psi, -\partial_x \psi, u_z)$ the Coriolis force can be absorbed in the pressure gradients $\frac{\partial \vec{u}}{\partial t} + (\vec{u} \cdot \vec{\nabla})\vec{u} = -\frac{1}{\rho}\vec{\nabla}\tilde{p} + \nu\Delta\vec{u}$ with $\tilde{p} = p + 2\rho\Omega\psi$ **2D 3C flow** \rightarrow The dynamics is no more affected by rotation

\rightarrow "Turbulence affected by rotation " is neither exactly 2D, nor at *Ro* = 0

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Navier-Stokes equation in a rotating frame

$$\frac{\partial \vec{u}}{\partial t} + (\vec{u} \cdot \vec{\nabla})\vec{u} = -\frac{1}{\rho}\vec{\nabla}p - 2\vec{\Omega} \times \vec{u} + \nu t\vec{u}$$

Restoring action of the Coriolis force

 $\frac{\partial \vec{u}}{\partial t} = -2\vec{\Omega} \times \vec{u}$

For velocities in the plane $\perp \overline{\Omega}$

$$\widehat{\Omega} = \Omega \overline{e_z}$$

Anticyclonic circular translation at frequency $\sigma = 2\Omega$

Inertial waves in fluids under rotation



$$\begin{cases} \frac{\partial u_r}{\partial t} = 2\Omega\cos(\theta) \, u_y \\ \frac{\partial u_y}{\partial t} = -2\Omega\cos(\theta) u_r \end{cases}$$

Anticyclonic circular translation of the plane tilted of θ at the frequency $\sigma = 2\Omega \cos(\theta)$

Dispersion relation
$$\frac{\sigma}{2\Omega} = \cos(\theta) = Ro_t$$

→ Propagation along
$$\theta = a\cos\frac{\sigma}{2\Omega}$$



Cortet, Lamriben & Moisy, POF (2010)



Coriolis force

$$\vec{F_c} = -2\vec{\Omega} \times \vec{u} = 2\Omega u_y \cos(\theta) \vec{e_r} - 2\Omega u_r \cos(\theta) \vec{e_y} - 2\Omega u_y \sin(\theta) \vec{e_\theta}$$

$$\frac{1}{\rho}\vec{\nabla}p = 2\Omega u_y \sin(\theta)\vec{e_{\theta}}$$

«
$$u_0\cos(\varphi(\vec{x},t))$$
»

$$\vec{\nabla}\varphi = \rho \frac{u_0}{p_0} 2\Omega \sin(\theta) \overrightarrow{e_{\theta}} \rightarrow \vec{\nabla}\varphi \perp \vec{u}$$

For a monochromatic wave $\vec{k} = \binom{k}{m} = \vec{\nabla}\varphi$

 $\rightarrow \frac{m}{\kappa} = \frac{\sigma}{2\Omega}$

and
$$\kappa = \sqrt{k^2 + m^2}$$
 arbitrary

Component along $\overrightarrow{e_{\theta}}$ taken by the pressure gradients





Gostiaux, Didelle, Mercier & Dauxois, Exp. in Fluids (2007) Mercier, Martinand, Mathur, Gostiaux, Peacock & Dauxois, JFM (2010)

UNIVERSITÉ

PARIS

construire l'avenir

Rotating platform with PIV on board



Excitation of a plane inertial wave

Bordes, Moisy, Dauxois & Cortet, POF (2012)



Dispersion relation
$$\frac{\sigma}{2\Omega} = \frac{m}{\kappa} = \cos(\theta)$$

- → Propagation along $\theta = a\cos\frac{\sigma}{2\Omega}$
- \rightarrow No link between frequency σ and wavelength
- \rightarrow Phase velocity normal to group velocity



Growth of two subharmonic waves

Temporal Fourier filtering





The subharmonic waves are plane waves

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New Challenges in Turbulence Research III

Resonance condition for a triad of plane waves

 $\sigma_1 + \sigma_2 + \sigma_3 = 0$

$$k_1 + k_2 + k_3 = 0$$

σ_0/f	$(\sigma_1 + \sigma_2)/f$	σ_1/f	σ_2/f
0.64	0.64	0.19	0.45
0.71	0.71	0.21	0.50
0.84	0.84	0.25	0.59
0.91	0.94	0.27	0.67
0.95	0.97	0.29	0.68
0.98	0.98	0.32	0.66
0.99	1.00	0.34	0.66



Resonance condition + dispersion relation

$$\begin{cases} \sigma_{1} + \sigma_{2} + \sigma_{3} = 0 \\ \mathbf{k}_{1} + \mathbf{k}_{2} + \mathbf{k}_{3} = \mathbf{0} \end{cases} + \frac{\sigma}{2\Omega} = s \frac{m}{\kappa} \quad \text{avec } \mathbf{k} = \binom{k}{m} \quad \text{et } s = \pm 1 \\ \implies s_{0} \frac{m_{0}}{\sqrt{k_{0}^{2} + m_{0}^{2}}} + s_{1} \frac{m_{1}}{\sqrt{k_{1}^{2} + m_{1}^{2}}} + \dots \\ - s_{2} \frac{m_{0} + m_{1}}{\sqrt{(k_{0} + k_{1})^{2} + (m_{0} + m_{1})^{2}}} = 0. \\$$

Decomposition of Navier-Stokes equation in helical modes

Cambon & Jacquin, JFM (1989) Waleffe, POF (1992)

Helical mode
$$\mathbf{u}(\mathbf{x},t) = A(t) \mathbf{h}_{\mathbf{s}} e^{i(\mathbf{k} \cdot \mathbf{x} - \sigma_s t)} + c.c.$$

with
$$\mathbf{h}_{s_k}(\mathbf{k}) = \frac{\mathbf{k}}{|\mathbf{k}|} \times \frac{\mathbf{k} \times \mathbf{e}_{\mathbf{z}}}{|\mathbf{k} \times \mathbf{e}_{\mathbf{z}}|} + is_k \frac{\mathbf{k} \times \mathbf{e}_{\mathbf{z}}}{|\mathbf{k} \times \mathbf{e}_{\mathbf{z}}|}$$

 \rightarrow Consider a velocity field made of an assembly of helical modes

$$\mathbf{u} = \sum_{\mathbf{k}} A_{\mathbf{k}}(\mathbf{k}, t) \, \mathbf{h}_s(\mathbf{k}) e^{i(\mathbf{k} \cdot \mathbf{x} - \sigma t)}$$

Navier-Stokes equation becomes

$$\left(\frac{\partial}{\partial t} + \nu\kappa^2\right) A_{\mathbf{k}} = \frac{1}{2} \sum_{\mathbf{p},\mathbf{q}} C_{\mathbf{k}\mathbf{p}\mathbf{q}} A_{\mathbf{p}}^* A_{\mathbf{q}}^* e^{i(\sigma_k + \sigma_p + \sigma_q)t}$$

with
$$C_k = C_{\mathbf{kpq}} = [s_q \kappa_q - s_p \kappa_p] (\mathbf{h}_{s_p}^* \times \mathbf{h}_{s_q}^*) \cdot \mathbf{h}_{s_k}^* / 2$$

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A helical mode = a plane inertial wave

Under rotation, if we take \mathbf{e}_{z} along $\mathbf{\Omega}$,

$$\mathbf{u}(\mathbf{x},t) = A(t) \mathbf{h}_{\mathbf{s}} e^{i(\mathbf{k} \cdot \mathbf{x} - \sigma_s t)} + \text{c.c.}$$





describes a plane inertial wave

Restricting to 3 inertial waves with initially $A_0 \neq 0$ et $A_{1,2} = 0$ $\left(\frac{\partial}{\partial t} + \nu \kappa^2\right) A_{\mathbf{k}} = C_{\mathbf{k}} A_{\mathbf{p}}^* A_{\mathbf{q}}^* e^{i(\sigma_{\mathbf{k}} + \sigma_{\mathbf{p}} + \sigma_{\mathbf{q}})t}$

> (i) The interaction (**k**, **p**, **q**) possible for **k** + **p** + **q** = **0** (ii) For $A_{1,2}(t)$ to grow, $\sigma_1 + \sigma_2 + \sigma_3 = 0$

$$\begin{cases} \frac{dA_1}{dt} = C_1 A_0^* A_2^* - \nu \kappa_1^2 A_1 \\ \frac{dA_2}{dt} = C_2 A_0^* A_1^* - \nu \kappa_2^2 A_2 \end{cases}$$

$$A_{1,2}(t) = B_{1,2} \left(e^{\gamma_+ t} - e^{\gamma_- t} \right)$$

with the growth rate

$$\gamma_{\pm} = -\frac{\nu}{2}(\kappa_1^2 + \kappa_2^2) \pm \sqrt{\frac{\nu^2}{4}(\kappa_1^2 - \kappa_2^2)^2 + C_1 C_2 |A_0|^2}$$



Energy transfers toward modes with more horizontal wavevectors



→ Two-dimensionalisation mecanism for weakly non-linear interaction

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$$\left(\frac{\partial}{\partial t} + \nu\kappa^2\right) A_{\mathbf{k}} = C_{\mathbf{k}} A_{\mathbf{p}}^* A_{\mathbf{q}}^* e^{i(\sigma_{\mathbf{k}} + \sigma_{\mathbf{p}} + \sigma_{\mathbf{q}})t}$$

(i) The interaction (**k**, **p**, **q**) possible for **k** + **p** + **q** = **0** (ii) For $A_{1,2}(t)$ to grow, $\sigma_1 + \sigma_2 + \sigma_3 = 0$

$$= \begin{cases} \frac{dA_1}{dt} = C_1 A_0^* A_2^* - \nu \kappa_1^2 A_1 \\ \frac{dA_2}{dt} = C_2 A_0^* A_1^* - \nu \kappa_2^2 A_2 \end{cases}$$

This framework makes sense if one have a separation of timescales

 $\gamma \ll \sigma_{1,2,3}$

Since
$$\gamma \sim \kappa A \rightarrow Ro \ll Ro_t$$

Moreover, in order to have a good phase mixing, the resonance needs to be precise to

$$\sigma_1 + \sigma_2 + \sigma_3 \ll \gamma = \Omega Ro$$

→ When *Ro* \land more non-resonant triads are allowed to exchange energy

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Decaying grid turbulence under rotation

with F. Moisy C. Lamriben



Lamriben, Cortet & Moisy, PRL (2011)





Camera

Turbulence decay, Ro(t)



Turbulence decay, Energy dissipation rate



Turbulence decay, Energy dissipation rate



$$\varepsilon = \frac{u_x'^2}{\tau_{tr}} = \nu_T \frac{u_x'^2}{L^2}$$

• without rotation $\varepsilon_{3D} = \frac{u_x^{\prime 3}}{L}$

• with rotation
$$\varepsilon_{rot} = \frac{u_x^{\prime 3}}{L_1} Ro = \varepsilon_{3D} Ro$$

 ε is not affected by rotation

→ Large scales are apparently not affected by rotation

Which scales affected?



Two points velocity correlation = « Energy distribution in the space of scales »

$$R(\vec{r},t) = \langle \vec{u}(\vec{x}+\vec{r}) \cdot \vec{u}(\vec{x}) \rangle$$

where $\langle \rangle$ stands for spatial (over \vec{x}) and ensemble average



 \vec{r} $\vec{u}(\vec{x} + \vec{r})$ $\vec{u}(\vec{x})$

 \rightarrow Isotropy \rightarrow Toward an anisotropic 2D state

How this anisotropy of $R(\vec{r}, t)$ is built by energy transfers? How it is distributed among scales?

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New Challenges in Turbulence Research III

Kármán-Howarth-Monin equation

Conservation of $R(\vec{r}, t)$ in the space of scales

$$< \mathrm{NS}(u_{i}(\vec{x} + \vec{r}))\mathrm{NS}(u_{i}(\vec{x})) > \rightarrow \frac{1}{2}\frac{\partial R}{\partial t} = \frac{1}{4}\vec{\nabla}\cdot\vec{F} + \nu\Delta R$$
Rotation has
disappeared

$$\prod(\vec{r},t) - D(\vec{r},t)$$
Recently distribution w

$$\vec{F}(\vec{r},t) = \langle (\delta\vec{u})^{2}\delta\vec{u} \rangle$$
Recently w

with $\delta \vec{u}(\vec{x}, \vec{r}) = \vec{u}(\vec{x} + \vec{r}) - \vec{u}(\vec{x})$ the velocity increment





KHM equation with rotation

$$\frac{1}{2}\frac{\partial R}{\partial t} = \frac{1}{4}\vec{\nabla}\cdot\vec{F} + \nu\Delta R$$
$$\underbrace{\prod(\vec{r},t)}_{\Pi(\vec{r},t)} - D(\vec{r},t)$$







- → Strong anisotropy of the energy transfers $\Pi(\vec{r}, t)$ between spatial scales
- → \vec{F} remains nearly radial but is dependent on the orientation of \vec{r}

Scale dependent anisotropy of energy transfers



The anisotropy of energy transfers $\Pi(\vec{r}, t)$ is stronger at small scales, i.e. decreases with r

 \rightarrow Consistent with the independence of *ε* with Ω.

Scale dependent anisotropy of energy distribution



Scale dependent anisotropy of energy distribution

Zeman, POF (1994)



- → The Zeman scale r_{Ω} is maybe a good description of the scale r_{aniso} .
- → But, the Zeman Rossby number seems not relevant for $r > r_{\Omega}$.

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The scale of maximum anisotropy



 $t V_g/M$

$$\frac{1}{2}\frac{\partial R}{\partial t} = \frac{1}{4}\vec{\nabla}\cdot\vec{F} + \nu\Delta R$$

$$\prod(\vec{r},t) - D(\vec{r},t)$$

Energy transfers $\Pi(\vec{r}, t)$ are more anisotropic at small scales but for $r < \eta$ they are no more at play $\rightarrow r_{aniso} \sim \eta$

A new Rossby number above the Zeman scale ?

→ For $r < r_{\Omega}$, the flow is not affected by rotation.

→ For $r > r_{\Omega}$, the Zeeman Rossby number $R_0(r) \sim \left(\frac{r_{\Omega}}{r}\right)^{\frac{2}{3}}$ fails to interpret data.

However, it is based on 3D isotropic scalings!!

 \rightarrow One needs to build a Rossby number increasing with *r*, in which case:

The larger the scale, the larger the Rossby number, the less affected by rotation

r



One can note that scale dependant Rossby numbers $R_0(r) \sim \frac{\sigma(r)}{2\Omega}$ associate a frequency $\sigma(r)$ with a scale r, which is highly incompatible with the physics of inertial waves!

Two-dimensionalisation of turbulence by rotation

Mininni, Rosenberg & Pouquet, JFM (2012)

DNS of forced turbulence under rotation with helical forcing

 $Re = 3 \times 10^4$ and Ro = 0.07

Delache, Cambon & Godeferd, POF (2014)

DNS of the decay of turbulence under rotation Initially isotropic and non-helical

 $Re = 5 \times 10^3$ and Ro = 0.02



Results are contrasted

The scale dependance of the anisotropy of rotating turbulence is still not fully understood. It probably depends on

- \rightarrow the helicity of the turbulence,
- \rightarrow the isotropy of the forcing or of the initial turbulence,
- \rightarrow the nature of the forcing (monochromatic or broadband).

→ Talk given by Pablo Mininni tomorrow 8h45 !

Rotating turbulence tends to be 2D but is not exactly.



What about the energy transfers in the horizontal plane $\perp \vec{\Omega}$?

An inverse cascade of energy?

Yarom, Vardi, Sharon, POF (2013)

Transient of a rotating turbulence experiment forced with an array of jets

Turbulence is both unstationnary and inhomogeneous Pouquet, Sen, Rosenberg, Mininni We are not sure when the the sen gray compart (2003)

Transistinof cates a tingph ES? forced at k_f





Kármán-Howarth-Monin equation

$$\frac{1}{4}\frac{\partial E(\vec{r})}{\partial t} = -\Pi(\vec{r}) + \frac{1}{2}\nu\Delta E - \varepsilon + P_r$$

For theory and numerics only

$$E(\vec{r},t) = \langle (\delta \vec{u})^2 \rangle = \langle (\vec{u})^2 \rangle - 2R$$

$$\Pi(\vec{r},t) = \frac{1}{4} \vec{\nabla} \cdot \langle (\delta \vec{u})^2 \delta \vec{u} \rangle$$

$$\varepsilon = \frac{1}{2} \nu \langle (\partial_i u_j + \partial_j u_i)^2 \rangle$$

$$P_r = \langle \delta \vec{u} \cdot \delta \vec{f} \rangle / 2 \sim P \text{ for } r > L_{\text{inj}}$$

$$\delta \vec{u}(\vec{x},\vec{r}) = \vec{u}(\vec{x}+\vec{r}) - \vec{u}(\vec{x})$$

In the isotropic case (2D or 3D), an angular average gives that $\rightarrow E(r)$ is the cumulated energy from 0 to r $\rightarrow \Pi(r_0)$ is the energy transferred from $r < r_0$ to $r > r_0$



3D turbulence

 $\Pi(r) = -\varepsilon < 0$

equivalent to $\langle (\delta u_L)^3 \rangle = -\frac{4}{5} \varepsilon r$ (Kolmogorov 1941, Monin & Yaglom 1975)

Forward cascade

2D turbulence

 $\Pi(r) = +P > 0$

equivalent to $\langle (\delta u_L)^3 \rangle = +\frac{3}{2}Pr$ (Kraichnan 1971, Bernard 1999, Lindborg 1999)

Inverse cascade

A stationary turbulence under rotation



A stationary turbulence under rotation

A vortex dipole flap generator



$$Re_f = \frac{\omega_f L_f^2}{v} \simeq 1000 \text{ and } Ro_f = \frac{\omega_f}{2\Omega} \simeq [0.03; 0.22]$$

The large scale vortex dipole produces and carries 3D turbulent fluctuations towards the center of the arena

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A stationary turbulence under rotation



Stereoscopic PIV

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New Challenges in Turbulence Research III

Flow in the vertical plane





 $Re \simeq 500$ $Ro \in [0.04 \ 0.15]$

Strong enhancement of the twodimensionality by rotation

Flow in the horizontal plane



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Energy flux density $\vec{F} = \langle \overline{(\delta \vec{u})^2 \delta \vec{u}} \rangle$



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$$\Pi(\overrightarrow{r_{\perp}}) = \frac{1}{4} \overrightarrow{\nabla} \cdot \langle \overline{(\delta \overrightarrow{u})^2 \delta \overrightarrow{u}} \rangle$$



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Energy flux $\Pi(r_{\perp})$ with an angular average

$$\Pi(r_{\perp}) = \langle \frac{1}{4} \vec{\nabla} \cdot \langle \overline{(\delta \vec{u})^2 \delta \vec{u}} \rangle \rangle_{\varphi}$$



Decomposition of the energy flux in vertical and horizontal energy

$$\partial_{t}\vec{u} + (\vec{u} \cdot \vec{\nabla})\vec{u} = \frac{1}{\rho}\vec{\nabla}p - 2\vec{\Omega} \times \vec{u} + \nu\vec{\Delta}\vec{u}$$

$$\begin{cases} \vec{u_{\perp}} = \begin{pmatrix} u_{x} \\ u_{y} \\ 0 \end{pmatrix} \\ u_{\parallel} = u_{z} \end{cases}$$

$$in the 2D3C limit$$

$$\partial_{t}\vec{u_{\perp}} + (\vec{u_{\perp}} \cdot \vec{\nabla_{\perp}})\vec{u_{\perp}} = \frac{1}{\rho}\vec{\nabla_{\perp}}\vec{p} + \nu\Delta_{\perp}\vec{u_{\perp}}$$

$$\partial_{t}u_{\parallel} + (\vec{u_{\perp}} \cdot \vec{\nabla_{\perp}})u_{\parallel} = \nu\Delta_{\perp}u_{\parallel}$$

$$u_{\parallel} = \text{passive scalar advected by the horizontal velocity}$$

$$Inverse cascade of horizontal energy$$

$$E_{\perp}(r_{\perp}) = \langle (\delta u_{\perp})^{2} \rangle$$

$$\Pi_{\perp}^{(1)}(r_{\perp}) = \frac{1}{4}\vec{\nabla_{\perp}} \cdot \langle (\vec{\delta}\vec{u_{\perp}})^{2}\vec{\delta}\vec{u_{\perp}} \rangle > 0$$

$$\Pi(r_{\perp}) = \Pi_{\perp}^{(\perp)}(r_{\perp}) + \Pi_{\perp}^{(\parallel)}(r_{\perp})$$

Decomposition of the energy flux in vertical and horizontal energy

$$\Pi_{\perp}^{\ (\perp)}(r_{\perp}) = \frac{1}{4} \overrightarrow{\overline{V_{\perp}}} \cdot \langle \overline{(\delta \vec{u}_{\perp})^2 \delta \overline{u_{\perp}}} \rangle > 0$$

Double cascade of energy

 $\Pi_{\perp}^{(\parallel)}(r_{\perp}) = \frac{1}{4} \overrightarrow{\nabla_{\perp}} \cdot \langle \overline{(\delta \vec{u}_{\parallel})^2 \delta \vec{u_{\perp}}} \rangle < 0$

3D isotropic forward energy transfer and/or stretching and folding of a passive scalar advected by $\overrightarrow{u_{\perp}}$ Pouquet, Sen, Rosenberg, Mininni & Baerenzung, Phys. Scripta (2013)

Rotating LES Energy injection at k_f Deusebio, Boffetta, Lindborg, Musacchio, ArXiv (2014)

DNS of turbulence under rotation Energy injection at k_f

with
$$A(\vec{r}) = -\frac{1}{4V_X} \oint_{S_X} \overline{(\delta \vec{u})^2 \vec{\tilde{u}}} \cdot d\overline{S_X}$$
 with $\tilde{\vec{u}} = \frac{\vec{u}(\vec{x}) + \vec{u}(\vec{x} + \vec{r})}{2}$
coarse-grained velocity
 $B(\vec{r}) = -\frac{1}{2V_X} \oint_{S_X} \overline{\delta p \delta \vec{u}} \cdot d\overline{S_X}$

$$D(\vec{r}) = \frac{\nu}{2} \vec{\nabla}^2 \langle \overline{(\delta \vec{u})^2} \rangle$$

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Kármán-Howarth-Monin equation in the horizontal plane

with
$$A = -\frac{1}{4V_X} \oint_{S_X} \overline{(\delta \vec{u})^2 \vec{\vec{u}}} \cdot d\vec{S_X}$$

 $B = -\frac{1}{2V_X} \oint_{S_X} \overline{\delta p \delta \vec{u}} \cdot d\vec{S_X}$??

$$D = \frac{\nu}{2} \vec{\nabla}^2 \langle \overline{(\delta \vec{u})^2} \rangle$$

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Kármán-Howarth-Monin equation in the horizontal plane

$$\varepsilon = -\Pi + D + \phi + B$$

 $\phi(r_{\perp}) = A + B$

with
$$A = -\frac{1}{4V_X} \oint_{S_X} \overline{(\delta \vec{u})^2 \vec{\hat{u}}} \cdot d \overrightarrow{S_X}$$

 $\frac{d\phi(r_\perp)}{dr_\perp} > 0 \rightarrow \text{energy injection}$
 $B = -\frac{1}{2V_X} \oint_{S_X} \overline{\delta p \delta \vec{u}} \cdot d \overrightarrow{S_X} ??$
 $\frac{d\phi(r_\perp)}{dr_\perp} < 0 \rightarrow \text{energy extraction}$
 $D = \frac{\nu}{2} \vec{\nabla}^2 \langle \overline{(\delta \vec{u})^2} \rangle$

A injects energy, ok!

But, where does the energy go at the end of the inverse cascade? Kármán-Howarth-Monin equation in the horizontal plane

$$arepsilon = -\Pi + D + \phi$$

 $\phi(r_{\perp}) = A + B$

$$\frac{d\phi(r_{\perp})}{dr_{\perp}} > 0 \rightarrow \text{energy injection}$$
$$\frac{d\phi(r_{\perp})}{dr_{\perp}} < 0 \rightarrow \text{energy extraction}$$

 dr_{\perp}

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Double cascade + injection of energy due to inhomogeneities

→ Direct measurement of energy transfers show that rotating turbulence can sustain a double cascade in the horizontal plane.

→ Inhomogeneities are necessary to explain the energy injection/extraction, even infinitely small.

There is need to consider the inhomogeneities in the statistics to account for the energy injection in an experimental stationary turbulence...

... whatever the degree of homogeneity of the rms velocity.

In our case, we are homogeneous in energy to within 3%.

An experimental homogeneous and stationary turbulence?

Lagrangian Exploration Module (12 propellers organised around a icosahedron)

Zimmermann et al. Rev Scient Inst (2010)

In the two presented experiments, we are either:

- homogeneous and decaying (no energy injection)
- inhomogeneous and stationary (energy injection)

These situations are not to avoid. They correspond to the reality of turbulence both in experiments and in nature. Trivial but often forgotten!

→ It is a necessary challenge to consider energy transfers in space and scales at the same time

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